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Article **Characteristics of Observed Electromagnetic Wave Ducts in** Tropical, Subtropical, and Middle Latitude Locations

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Abstract: Where and at what altitudes electromagnetic wave ducts within the atmosphere are likely to occur is important for a variety of communication and military applications. We examine 2 the modified refractivity profiles and wave duct characteristics derived from observed upper air soundings obtained over four years from 7 tropical and subtropical islands, and middle latitude sites 4 at 4 US coastal locations, 3 sites near the Great Lakes, and 4 US inland sites. Across all location types, 5 elevated ducts are more common than surface-based ducts and median duct thicknesses are ~ 100 m. There is a weak correlation between duct thickness and strength and essentially no correlation between duct strength and duct base height. EM ducts more frequently occur at the tropical and subtropical island locations (~60%) and middle latitude coastal locations (70%) as compared to less 9 than 30% of the time at the Great Lake and US inland sites. The tropical and subtropical island sites 10 are more likely than the other location types to have ducts at altitudes higher than 2 km which is 11 above the boundary layer height. 12

Keywords: modified refractivity; wave duct; trapping layer; upper air sounding

0. Introduction

Atmospheric refraction bends electromagnetic (EM) waves when the waves traverse 15 gradients in temperature and humidity [1-4]. In general, the refractive index in Earth's 16 atmosphere decreases with increasing height and as a consequence beam paths bend 17 downward relative to the surface compared to their path in a vacuum. Profiles of refractive 18 index permit calculation of EM beam paths. In some weather conditions, stable layers 19 can occur and yield adjacent atmospheric layers with distinct temperature and humidity 20 characteristics and sharp gradients in refractive index. In these circumstances, the beam 21 paths can be ducted wherein the waves are guided within a horizontal layer which allows 22 them to travel further than they would in normal conditions. Temperature inversions, 23 where temperature increases with increasing altitude, can yield trapping layers within 24 ducts if the gradients in temperature and humidity are strong enough. Ducting can be 25 caused by subsidence aloft, boundary layer inversions, or cooling near the surface such as 26 by nocturnal radiation inversions over land or warm dry air moving over a cooler body of 27 water. In calm, stable conditions over ocean, air in contact with the sea surface can become 28 saturated yielding ducting conditions ~ 10 m in thickness [4]. Evaporative downdrafts from 29 precipitating clouds can also yield trapping layers at any altitude below cloud base. Surface 30 cold pools originating from evaporative downdrafts will spread laterally so trapping 31 layers associated with them will vary in height with time. Globally, the highest ducting 32 probabilities are found in the Arabian Sea and in marine stratocumulus conditions in the 33 subtropics [5]. 34

Previous work has extensively addressed the theory of EM refraction [1,2,4,6]. The characteristics of ducts have been the focus of many studies using both observations and

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Figure 1. Locations of upper air sounding data used for refractivity profile EM duct analysis. Location type is distinguished by marker type. Color indicates groupings of a location type.

Table 1. Upper air sounding sample sizes and total duct counts for ducts with >40 m thickness and M > 1.7 at 18 locations over 4 years (2019-2022) by location type.

Location Type:	Total Soundings:	# of soundings with \geq 1 duct:	Percent of soundings with \geq 1 duct:	Duct Count:
Island	19,524	11,538	59.1%	32,352
Coastal	8,850	6,338	71.6%	12,938
Lake	8,871	2,367	26.7%	3,508
Inland	11,992	3,565	29.7%	5,403

modeling [e.g. 7–14]. Other work has utilized inversion methods which estimate refractivity 37 profiles from the measured signal and wave propagation models [e.g. 15]. In this paper, we 38 address the prevalence and characteristics of observed ducts both at the surface and aloft 39 using a high vertical resolution (\sim 5 m) upper air sounding data set. These high resolution 40 soundings provide new details on shallow ducting layers in the atmosphere that are not 41 possible to resolve with lower vertical resolution observations or model output [e.g. 5]. 42 Information on the geography and altitudes of frequent refractivity conditions conducive 43 for ducting is useful for navigation, communication, weather radar, as well as defensive 44 and offensive military applications [6,11]. 45

1. Materials and Methods

We use upper air soundings with a native vertical resolution of ~ 5 m from selected sites in the United States, its territories, and several Pacific islands. Data are from the period 1 January 2019 to 31 December 2022. In total, atmospheric profiles were analyzed for 49,239 upper air soundings, 23,806 of which contained one or more ducts (Table 1). We analyze profiles from 7 island sites, 4 coastal sites, 3 sites around the Great Lakes, and 4 inland sites (Fig. 1, Table 2). These categories represent varying geographic settings which influence atmospheric properties and ducting behavior. At coastal locations, offshore flow and alongshore flow can yield surface-based ducts and onshore flow can yield both surface-based and ducts aloft [9]. Previous work has not examined observed duct characteristics at multiple tropical island locations. While the lowest sounding levels are island influenced, once the sounding is a few km downwind of the island, the conditions are more representative of open ocean.

The sounding data are archived by the National Centers for Environmental Information (NCEI) in Binary Universal Form for the Representation of meteorological data (BUFR)

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Table 2. Upper air sounding locations subset into corresponding location type: tropical and subtropical island, US coastal, Great Lake, and US inland respectively. Latitude and longitude coordinates and local times at 00 and 12 UTC are listed for each location. For sites that participate in daylight savings, the first local time listed is daylight standard time and the second is daylight savings time.

Location Name	Latitude (°)	Longitude (°)	Local time at 00 UTC	Local time at 12 UTC		
Island						
American Samoa	-14.33	-170.71	1300	0100		
Chuuk	7.46	151.84	1000	2200		
Guam	13.48	144.80	1000	2200		
Marshall Islands	7.06	171.27	1200	0000		
Micronesia	6.99	158.21	1100	2300		
Puerto Rico	18.22	-66.59	2000	0800		
Yap	9.50	138.08	1000	2200		
		Coasta	ıl			
Oakland, CA	37.80	-122.27	1600/1700	0400/0500		
Newport, NC	34.79	-76.86	1900/2000	0700/0800		
Quilayute, WA	47.94	-124.54	1600/1700	0400/0500		
Tampa, FL	27.95	-82.46	1900/2000	0700/0800		
Lake						
Buffalo, NY	42.89	-78.88	1900/2000	0700/0800		
Gaylord, MI	45.03	-84.67	1900/2000	0700/0800		
Green Bay, WI	44.51	-88.01	1800/1900	0600/0700		
Inland						
Caribou, ME	46.86	-68.00	1900/2000	0700/0800		
Fort Worth, TX	32.76	-97.33	1800/1900	0600/0700		
Minneapolis, MN	44.98	-93.27	1800/1900	0600/0700		
Nashville, TN	36.16	-86.78	1800/1900	0600/0700		



Figure 2. Idealized schematic of electromagnetic wave duct components as a function of modified refractivity and height. Adapted from [9].

format [16,17]. Operational upper air soundings are launched at \sim 11 UTC and \sim 23 UTC to achieve mid troposphere altitudes at 0 and 12 UTC each day. The data set we use spans 10 time zones, yielding differences in local times among the sites (Table 2). We do not have adequate temporal sampling to analyze diurnal cycle variations.

The observed upper air sounding profiles at ~ 5 m native resolution are linearly interpolated to 20 m vertical layers and then input to the calculation of modified refractivity (*M*). Modified refractivity is a function of temperature, water vapor, pressure and the curvature of the Earth. The advantage of modified refractivity over refractivity is that all negative *M* vertical gradients are associated with trapping layers which simplifies duct identification[4,15]. Modified refractivity (*M*) is determined using the following equation:

$$M = \frac{77.6}{T} \left(P + \frac{4810e}{T} \right) + \frac{z}{10^{-6}R_e} \tag{1}$$

where *P* is pressure (mb), *T* is temperature (K), *e* is vapor pressure (mb), *z* is height (m) and R_e is the radius of the Earth (m) [2,7]. Modified refractivity values are calculated for each height level in each sounding.

Figure 2 illustrates the key components and characteristics of a wave duct as a function 74 of modified refractivity and altitude. A trapping layer is characterized by a decrease in 75 modified refractivity with increasing height [9]. The thickness of an elevated duct is the 76 distance between the local minimum in M above the trapping layer to the same value of M77 below the trapping layer. Surface-based ducts only have the trapping layer portion in 78 which case the duct thickness is defined as the trapping layer thickness. Observed examples 79 of modified refractivity profiles are annotated with trapping layer top and bottom and duct 80 base in Figure 3. The example from Guam at 1106 UTC on 31 March 2022 contains seven 81 ducts aloft (Fig. 3a). A surface duct along with three ducts aloft is shown in the example 82 from Wallops Island, VA, at 2300 UTC on 11 July 2022 (Fig. 3b). 83

Using the information on trapping layer top and bottom and duct base, we calculate duct strength as the difference between the local maxima in M at the trapping layer bottom and local minima in M at the trapping layer top. Two thresholds were applied to filter out very weak or very thin modified refractivity inversions. A duct was included in the analysis only if it met both the criteria of duct strength M > 1.7 and duct thickness > 40 m. By focusing on these features, this study aims to provide improved understanding of electromagnetic wave duct frequency of occurrence and variations across different environments.

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Figure 3. Modified refractivity profiles from a) Guam on 1106 UTC 31 March 2022 with multiple ducts aloft and b) Wallops, VA on 2300 UTC 11 July 2022 with multiple ducts aloft and a surface duct. Planetary boundary layer height based on Richardson number and on virtual potential temperature (Theta V) [18] differed in a) and are shown separately. In b) both estimates of the boundary layer height were at the same altitude.



Figure 4. Distribution of number of ducts in a given sounding for a) tropical and subtropical island locations, b) US inland, c) US coastal, and d) Great Lake locations. For island locations, 39 soundings had more than 13 ducts.

	10th	25th	Median	75th	90th	
Strength (M)						
Island	2.04	2.66	4.13	7.23	12.40	
Coastal	2.10	2.83	4.69	8.69	15.11	
Lake	2.03	2.69	4.23	7.46	12.42	
Inland	2.05	2.70	4.35	8.19	14.46	
Thickness (m)						
Island	52	66	94	156	257	
Coastal	55	73	114	191	298	
Lake	54	72	104	164	238	
Inland	58	76	113	179	276	
Duct Base Height (m)						
Island	503	963	1781	2628	3589	
Coastal	55	324	874	1730	2720	
Lake	41	460	1128	1826	2599	
Inland	83	649	1320	2060	2867	
Duct Top Height (m)						
Island	627	1098	1935	2750	3689	
Coastal	219	512	1018	1887	2720	
Lake	177	579	1272	1969	2708	
Inland	228	810	1490	2198	2975	

Table 3. Median and 10th, 25th, 75th, 90th percentiles for duct characteristics strength, thickness, duct base height, and duct top height for each location type.

2. Results

Most atmospheric profiles at coastal (72%) and island (59%) locations have one or 92 more ducts (Table 1). In comparison, the lake and inland profiles had one or more ducts 93 less than 30% of the time. Multiple ducts per profile are more likely in island locations 94 (Fig. 4). Statistics for duct characteristics by location type are presented in Table 3. Median 95 duct strengths (\sim 4.1 to 4.7 M) and thicknesses (\sim 100 m) are similar among the location 96 types. The distributions of duct base altitudes show notable differences among location 97 types, with the median value for islands (1781 m) about 1 km higher in altitude than for 98 coastal (874 m). As a consequence, median duct top altitudes are also about 1 km higher qq for island as compared to coastal sites. Great Lake and inland locations have median duct 100 base altitudes, 1128 m and 1320 m respectively, at intermediate values between the island 101 and coastal values. The increased height of ducts between US coastal versus subtropical 102 and tropical islands is consistent with the increasing height of the inversion-topped marine 103 boundary layer documented along ship transects that traversed from marine stratocumulus 104 to trade cumulus conditions between Southern California and Hawaii [12]. 105

The interrelationships among the heights, thicknesses, and strength of the ducts 106 observed between the locations are illustrated in Figures 5, 6, and 7. Most ducts are not 107 surface-based (Fig. 5). 75% have tops within the first 3 km of the surface and 75% have 108 thicknesses of 200 m or less (Fig. 6, Table 3). As duct thickness increases > 200 m, the range 109 of duct strengths tends to broaden to include higher value outliers. While thin ducts tend 110 to be weak (M < 10) the high prevalence of weak, thick ducts yields linear correlations that 111 explain less than half the variance between duct strength and and thickness (Fig. 6). There 112 is no meaningful linear correlation between duct strength and height (Fig. 7). 113

a) Island

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Height (km)

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Figure 5. EM duct thickness (vertical length of line) and strength (color-coded) sorted by duct top height for a) tropical and subtropical islands, b) US inland, c) US coastal, and d) Great Lake locations. For island locations, thirteen ducts had duct top heights exceeding 7 km.

The range of surface-based duct strengths is similar to that for elevated ducts (Fig. 114 7). There are a few outliers with M > 30 in each location category. The surface-based 115 ducts represent a combination of evaporative downdrafts reaching the surface, nocturnal 116 radiation cooling, and sea breezes in coastal, lake, and island locations. Nocturnal radiation 117 inversions are more likely to form on calm, clear nights at inland locations than near 118 bodies of water. The onshore movement of a low layer of cooler ocean air in the afternoon 119 associated with the sea breeze (or a lake breeze) would be conducive to ducting. 120

The frequency of occurrence of soundings with ducts in different atmospheric layers 121 (surface, between surface and 2 km, and \geq 2 km altitude) is tabulated in Table 4. If a 122 particular sounding had a surface-based duct and one duct above 2 km it would be counted 123 in each of those categories. The island locations stand out with many more soundings 124 containing ducts above 2 km altitude (61%) compared to the other locations (< 31%) which 125 is discussed further in Section 3. 126

Duct characteristics from US coastal measurements are usually not representative of 127 tropical open ocean conditions. Compared to coastal ducts, the distribution of subtropical 128 and tropical islands ducts heights are shifted to higher altitudes (Fig. 5 and 7, Table 3). 129 Island locations are more likely to have stronger ducts (M > 25) at altitudes above 1 km 130 than coastal locations (Fig. 7). The frequency of surface-based ducts is lower on islands 131 (2%) compared to coastal locations (10%) (Table 4). 132

3. Discussion

For island locations, while surface-based ducts, and ducts associated with inversions 134 near boundary layer tops were expected, the high prevalence of ducts with bases above 135 boundary layer height (duct base > 2 km altitude) was not expected (Fig. 5, 7, Table 4). Ducts 136 are stable layers so a key question is what are the likely mechanisms producing stable layers 137



Figure 6. Scatter density plots of duct thickness (m) vs. strength (M) for a) tropical and subtropical islands, b) US inland, c) US coastal, and d) Great Lake locations. Linear regression line (red), corresponding equation, and the coefficient of determination (r^2) are displayed on each subplot. Shading indicates number of samples.



Figure 7. Scatter density plots of duct base height (m) vs. strength (M) for a) tropical and subtropical islands, b) US inland, c) US coastal, and d) Great Lake locations. Shading indicates number of samples.

Table 4. Counts and percentages of duct occurrences in different layers of the atmosphere based on location type. In most circumstances, altitudes > 2 km would be above the boundary layer height [18]. Percentages are relative to total number of soundings with one or more ducts for a location type (Table 1).

Criteria	Island	Coastal	Lake	Inland
Includes a surface-based duct	281 (2.4%)	631 (10.0%)	283 (12.0%)	393 (11.0%)
Includes \geq 1 duct with base >surface and <2 km	8724 (75.6%)	5233 (82.6%)	1829 (77.3%)	2600 (72.9%)
Includes \geq 3 ducts with base >surface and <2 km	1475 (12.8%)	1125 (17.8%)	140 (5.9%)	189 (5.3%)
$\begin{array}{l} \text{Includes} \geq 1 \text{ duct} \geq 2 \text{ km} \\ \text{ altitude} \end{array}$	7051 (61.1%)	1855 (29.3%)	602 (25.4%)	1120 (31.4%)
$\frac{1}{1} Includes \geq 3 ducts \geq 2 km \\altitude$	1637 (14.2%)	132 (2.1%)	25 (1.1%)	51 (1.4%)

 \geq 2 km altitude in tropical and subtropical oceanic settings? Subtropical marine regions 138 have persistent large-scale subsidence associated with the downward branch of the Hadley 139 circulation. The subsidence manifests as temperature inversions and humidity gradients 140 yielding trade wind cumulus clouds over warmer oceans and stratocumulus clouds over 141 cooler oceans [19,20]. However, large scale subsidence would not readily explain multiple 142 ducts aloft at different altitudes in the same sounding as illustrated in Figure 3a. In the 24 143 hours prior to the sounding, Guam hourly METARS reported up to three distinct cloud 144 layers with bases ranging between 600-2700 m. At any one time in a marine cumulus 145 cloud field, clouds are forming and dissipating. Some small cumulus clouds produce 146 precipitation that reaches the surface but most do not. Some cumulus yield virga which 147 will cool and moisten the air just below cloud base. Dry air entrainment dissipates clouds 148 and moistens and cools the immediate vicinity [e.g. 21]. Layer moistening by cumulus 149 cloud dissipation is implicated in the multi-week transition between suppressed (dry) 150 and active (wet) phases of the Madden-Julian oscillation [e.g. 22]. Over hourly to daily 151 time scales, layer moistening by cloud detrainment and virga may have implications for 152 the creation of \sim 100 m thick stable layers and potentially ducts if the moisture gradient 153 persists. Determining the physical mechanisms yielding multiple ducts aloft in these 154 settings requires in depth study with more data. 155

4. Conclusions

Our analysis of more than 49,000 modified refractivity profiles derived from 20 m vertical resolution upper air sounding data complements previous studies on EM duct characteristics based on coarser vertical resolution modeling and observations. By examining a large data set from geographically diverse sites we are able to discern similarities and differences among ducts in different environments. Key findings from observed profiles of modified refractivity including ducts at least 40 m in thickness and with strengths $M \ge 1.7$ are:

- In all location types, elevated ducts are more common than surface-based ducts.
- Median values of duct strengths (M between 4.1 and 4.7) and thicknesses (~100 m are similar across location types.
- Duct strength tends to increase with increasing duct thickness but this relationship explains less than half of the variance.
- Duct strength and duct base height are not correlated.
- Profiles with one or more ducts are common at tropical and subtropical island (~60%) 170 and US coastal locations (~70%) and occur less than 30% of the time at the Great Lakes 171 and US inland sites. 172

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- Notable differences between ducts at tropical and subtropical islands versus US coastal locations include islands having higher median duct base altitudes, higher frequency of stronger ducts at altitudes > 1 km, and lower frequency of of surface-based ducts.
- Tropical and subtropical island locations often exhibit one or more elevated ducts above 2 km altitude in a single profile, a phenomena requiring further investigation.

In the future, the duct inventory we have posted on a public archive can be used as input to radar propagation models to determine the impacts of the observed ducts aloft and at the surface on beam paths for different EM applications.

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